LA-UR-87-1370

LA-UR--87-1370

DE87 009004

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An application of accelerator mass spectrometry in geochronology

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SUBMITTED TO: Fourth International Symposium on Accelerator Mass Spectrometry

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MEASUREMENT OF COSMOGENIC ³⁶C1/C1 IN YOUNG VOLCANIC ROCKS: AN APPLICATION OF ACCELERATOR MASS SPECTROMETRY IN GEOCHRONOLOGY

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Abstract

We have measured 36 Cl/Cl ratios in a number of young volcanic rocks in order to test the feasibility of using 36 Cl buildup as a geochronometer for materials less than about 700,000 years old. All of the analyzed rocks have been dated independently using K-Ar or other radiometric dating methods and have exposure histories that are known or can be reasonably assumed. Measured 36 Cl/Cl ratios in these rocks are in good agreement with the calculated in-situ 36 Cl buildup curve. These analyses indicate that AMS measurement of 36 Cl buildup in young rocks is a potentially powerful new method for dating materials that had previously been undatable, and as such will have broad applications in volcanology, tectonics, geophysics, and Quaternary research.

Introduction

The dating of young volcanic rocks is a geological problem that has not been satisfactorily resolved. A number of methods have been used for dating relatively old rocks, but techniques for rocks less than 500,000 years old have not been as successful. Rocks in the 10,000-100,000 year range have been analysed using a number of techniques, including K-Ar (Dalrymple, 1967), fission-track (Naeser, 1979), radiocarbon (Lockwood and Lipman, 1980), 40 Ar/ 39 Ar (di Brozolo and others, 1981), and uranium-series disequilibrium (Allegre, 1968) but in many cases these have yielded equivocal results. Moreover, these methods are not applicable to all young volcanic rocks, due to restrictions on the types of materials suitable for dating.

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The commonly used radiometric dating techniques mentioned above are of two general types. One, of which K-Ar is an example, utilizes the buildup of daughter products of primordial radionuclides; the other, exemplified by standard ¹⁴C dating, uses the decay of cosmogenic isotopes that are produced in the atmosphere and then incorporated into terrestrial reservoirs. This study investigates a third approach, the measurement of cosmogenic radionuclides in geological materials exposed to cosmic rays at the surface of the earth. While the first two techniques measure the time since an object to be dated became a closed geochemical system, the third method measures the time of exposure at the surface. Therefore, this method has the potential for dating features which are currently undatable, such as geomorphic surfaces, as well as determining the exposure age of volcanic rocks.

The use of light, rare isotopes such as 36 Cl has been considered for dating young materials for some time (Davis and Schaeffer, 1955; Lal and Peters, 1962). However, levels of these isotopes on the earth are quite low, and it has only been since the development of accelerator mass spectrometry (AMS) that an analytical method sensitive enough for practical analysis has been available. Recently, 36 Cl studies have been applied to groundwater (Bentley and others, 1984) lacustrine sediments (Phillips and others, 1983) and geothermal systems (Phillips and others, 1984), but have not been applied to the dating of igneous rocks.

Cosmogenic 36Cl Buildup Dating

 36 Cl has several characteristics which made it a prime candidate for buildup dating. It builds up to measurable levels relatively quickly since the product element (chlorine) is present in most rocks only in trace quantities; 36 Cl activities produced by nuclear processes within rocks, due to the decay of uranium and thorium, are much lower than the cosmogenic activity after only a few thousand years of exposure; and the mobile, hydrophilic nature of chlorine aids in the separation of <u>in-situ</u> produced 36 Cl from that derived from an atmospheric (meteoric) source. Recent research (Kimura and others, 1983) has shown that even relatively immobile radionuclides can penetrate rocks in geologically short time periods. Since 36 Cl is more hydrophilic than either 10 Be or 26 Al – two other candidates for

use in buildup dating (Nishiizumi and others, 1986) - this feature is useful in cases where it is necessary to separate <u>in-situ</u> and meteoric radionuclides. Most 36 Cl in rocks is produced by thermal neutron activation of 35 Cl; a second method of production, important mainly in rocks with low chlorine content, is direct spallation of 36 Cl from 40 Ca and 39 K. Spallation from titanium and iron is insignificant at the earth's surface due to attenuation by the atmosphere of the highly energetic cosmic rays necessary for these reactions. Calculation of the spallation production rate is discussed in Yokoyama and others (1972) and representative values for spallation production are given in Yokoyama and others (1977). Negative muon capture by 40 Ca (Kubik and others, 1984) and neutron activation of 39 K are minor reactions that produce 36 Cl. Although potentially significant at depth, these reactions may be ignored in calculations for surface rocks.

Cosmogenic isotopes accumulate at a rate proportional to the exposure time of the material; at the same time, they decay at a rate proportional to their decay constants. The measured 36 Cl content of a rock (usually expressed as the ratio of the number of atoms of the 36 Cl to the number of atoms of stable chlorine) is thus a function of the production rate, decay rate, and initial element content of the rock (Bentley and others, 1985), given as:

$$(^{36}C1/C1)_{+} = ([ED(PR)]/\lambda^{35}C1)(1-exp(\lambda t)) + (^{36}C1/C1)_{0}exp(-(\lambda t))$$
 (1)

where

 λ = decay constant

t= time in years

PR= production rate of 36 Cl

 35 Cl= number of atoms of 35 Cl

E= factor accounting for elevation of sample site

D= factor accounting for depth of sample below surface

and the subscripts (m) and (0) refer to measured and initial ratios, respectively.

The E term corrects for the attentuation of cosmic rays in the atmosphere; tables for this correction are given in Yokoyama and others (1977). Lal and Peters (1967) provide D, the depth correction for subsurface samples.

Sampling

In order for Equation (1) to be applied to surface exposure dating in a straightforward manner, the material to be dated should have been completely shielded from cosmic-ray exposure until exposed suddenly at the surface. Volcanic rocks admirably meet these conditions, and there are enough very young rocks that have been dated by conventional methods to provide a sufficient number of calibration samples. Hence these are an ideal subject for a test of a cosmogenic isotope buildup chronometer. Our samples include young volcanic rocks that are fresh, relatively dense and impermeable to groundwater, located at high elevations that will provide maximum cosmic-ray flux, and have been dated by reliable methods with ages that are compatible with the observed geology of the sample site. Also, wherever possible, samples were collected through a vertical section at a given site in order to allow evaluation of the variation of cosmogenic isotope production rates with depth below the ground surface. In addition to the isotopic measurements, these samples were analysed petrographically and geochemically so that the calibration suite would be fully documented. The rocks collected from the several localities within the arid regions of New Mexico and southeastern California are our primary suite of calibration samples; those samples collected from the more humid regions of California and the Pacific Northwest comprise a secondary group.

Preparation and analysis

The samples were prepared by splitting to remove any surface rind and then crushed and sieved to <100 mesh. The ground samples were leached with deionized water to remove adsorbed meteoric chloride. After leaching, the rock powders were digested by total fusion with NaOH and dissolved in deionized water. Chloride was precipitated from this

solution as AgCl by addition of an excess of $AgNO_3$. Since ^{36}S is an interfering nuclide in AMS analysis of ^{36}Cl , the AgCl was carefully purified by repeated dissolution in NH_4OH and reprecipitation with HNO_3 . $^{26}Cl/Cl$ ratios were determined using the TAMS at the Nuclear Structure Research Laboratory at the University of Rochester (Elmore and others, 1979). Whole-rock major and minor element analyses were performed by x-ray fluorescence and neutron-activation analysis at the Los Alamos National Laboratory.

Results

Since the analysed rocks were collected at different locations and elevations and are of different chemical compositions, all measured 36 Cl/Cl ratios were normalized to sea-level cosmic-ray flux and an arbitrary reference composition. Normalized ratios are shown relative to the predicted buildup for the reference composition on Figure 1. The values were normalized by:

$$[(^{36}\text{C1/C1}) - (^{36}\text{C1/C1})_0] (^{C1}/\text{C1}_r) [K_20(PR_K)_r + CaO(PR_{Ca})_r + C1_r(PR_a)_r] / [K_20(PR_C)_r + CaO(PR_{Ca})_r + C1_r(PR_a)_r] / (ELD)$$
 (2)

Reference compositions were 165 ppm Cl, 2.35 weight per cent K_2O , and 5.00 weight percent CaO. Samples from depth were normalized using their density as measured by specific gravity and a half-attenuation length of 132 g/cm^2 (Nishiizumi and others, 1984). The production rates for 36 Cl from spallation of K and Ca were 2670 atoms/kg per 36 Cl atoms/kg per 36 Cl (Yokoyama and others, 1977). The reference production rate of 36 Cl by neutron activation of 35 Cl was 81 atoms/year per ppm Cl, based on a sea-level neutron flux of 36 neutrons/kg per year (Lal and Peters, 1967).

Discussion

Our results show that 36 Cl/Cl increases with sample age in a predictable manner. Several of the samples, mainly rocks of rhyolitic composition, have higher values than expected. Since our early work

indicated that these samples were not contaminated with meteoric "bomb" ³⁶Cl, these high values suggest that the cosmogenic nuclide production rates that we have been using may be in error. In comparison, the values for the near-surface basalts compare well with the predicted curve. As the rhyolites have a significantly higher K content than the basalts (4-5% K_20 vs. 1-2% K_20), it would appear that the production rate for ³⁶Cl from K could be higher than published values. We have been using rates extrapolated from data on the spallation production of ²²Na from Al (Yokoyama and others, 1977); it may be that the 39 K(n, \propto) reaction is more important at near-surface levels than previously realized. Also, several of the basalts sampled at depth have similarly higher ratios than expected, which suggests that the muon-induced production of ³⁶Ci from Ca may likewise be significant for these rocks. Indeed, Kurz (1986) reported muon-induced production of ³He to be the dominant mechanism for production of this isotope at depths below 170 g/cm^2 , and the same may be true for ^{36}Cl . However, it will be necessary to accumulate more data before these production rates can be fully ε aluated. In addition to our data for volcanic rocks, preliminary ³⁶Cl/Cl analysis of granitic boulders from Tioga-age (10,00-20,000 years b.p.) moraines from the Sierra Nevada (Phillips and Jannik, unpub. data) indicate that buildup dating will be applicable to this type of material as well.

Summary

 36 Cl analyses have been obtained for a number of well-dated young volcanic rocks as part of an evaluation of the use of this isotope as a dating tool. Although uncertainties remain in the interpretation of our data, it is apparent that there is a predictable huildup of 36 Cl with time, produced in-situ since the emplacement of these flows at the surface. We expect that, with refinement, 36 Cl buildup can be used to date surface-exposure ages for volcanic rocks and other surficial materials. The ability to directly date very young materials of a wide range of compositions, such as archeological artifacts, glacial moraines, geomorphic surface, and soils, as well as volcanic rocks, will constitute a powerful tool for many of the earth sciences.

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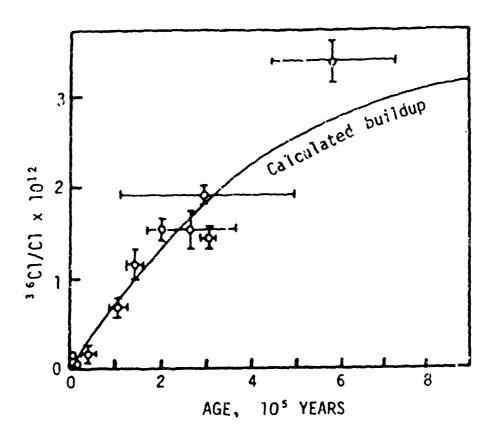


Figure 1: Normalized 36 Cl/Cl in calibration samples vs. published ages of sampled flows.